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Relations of Geomorphology to Stream Habitat and Trout Standing Stock in Small Rocky Mountain Streams

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Abstract.—Evidence that drainage basin morphology and trout standing stock are related through a functional link between geomorphic features and stream habitat quality is presented. Numerous significant univariate correlations were found between geomorphic variables, stream habitat variables, and trout standing stock in both high-elevation forest and low-elevation rangeland streams. Canonical correlations between geomorphic variables and stream habitat variables provided insight into the form of the functional link. Multiple-regression equations predicting trout standing stock were dominated by geomorphic variables. When geomorphic variables alone were incorporated into regression models they predicted trout standing stock as accurately as did stream habitat variables.

Methods for predicting standing stock of trout (species of *Salmo* and *Salvelinus*) in Rocky Mountain streams have focused mainly on stream habitat variables; little attention has been given to the possible influence of drainage basin geomorphology on stream habitat quality. Streams are known to reflect both the hydrology and biology of their watersheds (Platts 1979), but fish production may also be related to geomorphic processes in the drainage basin.

A few studies have attempted to relate geomorphic features of the watershed with salmonid standing stocks. Using geomorphic variables from Ziemer (1971) and Burton and Wesche (1974), Wesche et al. (1977) developed an index of habitat quality for cutthroat trout *Salmo clarki* in the Sierra Madre Range of Wyoming. Oswood and Barber (1982) combined measures of drainage basin geomorphology and stream habitat to predict salmonid standing stock in Alaskan streams, whereas Parsons et al. (1981) developed models incorporating geomorphic variables for salmonid streams in Oregon. However, these studies have not investigated the relation between geomorphic variables and stream habitat variables, nor have they examined the contribution of each type of variable when predicting salmonid standing stock in streams.

We demonstrate that measures of drainage ba-

sin geomorphology are related to both stream habitat features and trout standing stock. We also describe the ability of geomorphic and stream habitat variables to predict trout standing stock independently and in combination with each other.

Methods

Data were compiled for streams in the Colorado and Missouri river drainages within Wyoming from two sources: file data from the U.S. Bureau of Land Management and the University of Wyoming, and data gathered in the field during summer 1984. File data were accepted if three conditions were met: (1) stream habitat data and standing stock estimates were collected over the same reach within 1 month of each other; (2) a channel-stability evaluation had been conducted (Pfankuch 1975), and nine stream habitat variables measured (Table 1); and (3) a minimum two-pass depletion estimate of fish abundance had been made by the DeLury (1951) or Zippin (1958) methods.

Stream habitat variables.—In June, July, and August 1984, data were collected on small perennial streams (< 10 m average wetted width during summer low-flow) known to support trout. At least one pool-riffle sequence typical of the stream was included in each 75-m study reach. Within each reach, 10 cross-channel transects were established at 7.5-m intervals. Wetted stream width was measured perpendicular to flow at each transect and mean wetted stream width was then computed for the reach. Depth measurements were taken at

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TABLE 1.—Range of measured values for stream reaches assessed in Wyoming.

Variable	Forest streams (N=65)	Rangeland streams (N=26)
Trout standing crop (kg/hectare)	1.0–604.2	8.5–393.9
Stream measurements		
Channel stability score ^a	51–141	61–128
Average wetted reach width (m)	0.78–9.14	1.52–7.47
Average reach depth (m)	0.04–0.40	0.50–0.46
Average reach velocity (m/s)	0.06–0.81	0.07–0.74
Width : depth ratio	5.0–88.83	6.63–48.80
Bedrock–boulder substrate (%)	0–74	0–53
Rubble substrate (%)	0–70	0–67
Gravel substrate (%)	0–70	3–56
Silt–sand substrate (%)	0–46	6–77
Reach gradient (%)	1–9	1–4
Geomorphologic measurements		
Reach elevation (m)	2,097–3,158	1,329–2,245
Midrange basin elevation (m)	2,426–3,362	1,987–2,841
Stream order	1–5	2–6
Basin area (hectare)	95–39,290	1,348–48,918
Basin perimeter (ha)	5–96	23–173
Basin relief (m)	165–1,601	267–3,024
Compactness coefficient ^b	0.08–0.36	0.14–0.26
Stream length (km)	1.7–29.3	8.3–72.4
Relief ratio (m/km)	23.5–262.3	14.8–116.3
Channel slope (m/km)	8.4–116.1	10.1–70.0
Drainage density (km/km ²)	0.40–4.2	0.8–5.5

^a Low values indicate stability, high values erosive conditions.

^b Basin perimeter/2(3.14·basin area)^{1/2}.

points that were 0.25, 0.50, and 0.75 of the wetted stream width; the three depth measurements for each transect were summed and divided by four to compute mean transect depth. Platts et al. (1983) found this method of computing mean depth to have a 95% confidence interval about the mean of $\pm 8.2\%$. The mean depth for each of the 10 transects was averaged to obtain mean depth of the reach. Width : depth ratio was computed as the mean wetted width divided by the mean reach depth. At each point where depth was measured, the dominant substrate class was visually determined as either silt–sand (≤ 0.25 cm diameter), gravel (0.26–7.5 cm), rubble (7.6–30.0 cm), or bedrock–boulder (≥ 30.1 cm). The number of points at which each substrate class was found, divided by the total number of measurements, gave the proportion of each substrate class in the reach.

To estimate surface water velocity, we floated a pencil three times over a relatively straight, unobstructed subsection of the reach for about 20 s. Distance traveled and float duration were recorded. The mean subsection velocity (m/s) was computed and multiplied by 0.85 to adjust for above-average water velocity at the surface. Within this subsection, three equally spaced transects were established to determine average subsection width and depth. Stream discharge (m³/s) was cal-

culated as the mean cross-sectional area of the reach subsection multiplied by mean velocity in the subsection. For an assumed constant discharge through the reach, mean reach velocity was computed as the discharge divided by the mean cross-sectional area of the 10 reach transects. Reach gradient was estimated with a clinometer.

Channel stability was visually estimated following Pfankuch (1975). Fifteen stability indicators were rated numerically over an entire stream reach and summed to yield a reach score used in our data analyses. The score reflects the channel stability, with a low value indicative of a stable channel and a high score indicative of an erosive channel.

Geomorphologic variables.—Eleven geomorphologic variables were measured on 1:24,000 or, when not available, 1:62,500 scale topographic maps of the U.S. Geological Survey (Reston, Virginia) (Table 1). Each study reach was located on a topographic map and its drainage divide was drawn. Variables were measured as follows:

- (1) Study reach elevation: read directly from the map.
- (2) Midrange basin elevation: (highest elevation on the headwater divide + reach elevation)/2.

- (3) Stream order: determined by counting the stream channels shown in blue on topographic maps (Horton 1945, as modified by Strahler 1957).
- (4) Basin area: measured with a compensating polar planimeter (Horton 1945).
- (5) Basin perimeter: measured with a map measurer (Horton 1945).
- (6) Basin relief: highest elevation on the headwater divide minus the elevation of the reach (Schumm 1956).
- (7) Compactness coefficient: basin perimeter/2(3.14·basin area)^{0.5} (Parsons et al. 1981).
- (8) Stream length: measured by following the longest watercourse shown in blue on the map with a map measurer (Horton 1945).
- (9) Relief ratio: basin relief/stream length (Schumm 1956).
- (10) Channel slope: (elevation at 85% of stream length - elevation at 10% of stream length)/stream length between these two points (Craig and Rankl 1978).
- (11) Drainage density: length (km) of all stream channels shown in blue in a drainage basin/drainage area (km²) (Horton 1945).

Standing stock estimates.—Estimates of trout (brown trout *Salmo trutta*, rainbow trout *S. gairdneri*, brook trout *Salvelinus fontinalis*, and cutthroat trout) standing stock in each reach were made by the removal method (DeLury 1951). Each reach was blocked at the upper and lower ends with minnow seines and two or three depletion passes were made with a battery-powered backpack electroshocker. At the end of each pass, fish were weighed to the nearest gram and natural total length was measured to the nearest millimeter. Only data from trout 100 mm or longer were recorded. Estimates of trout abundance in each reach were computed with program CAPTURE (White et al. 1982). Model M(bh) was chosen because it allowed for variability in capture among animals and for behavioral responses to the first capture attempt.

Data analysis.—High-elevation coniferous forest watersheds were separated from lower-elevation rangeland watersheds. The boundary elevation between forest and rangeland streams approximately followed the low-elevation coniferous forest timberline in Wyoming: 2,287 m at 41°–41°60'N latitude; 2,135 m at 42°–42°60'; 1,982 m at 43°–43°60'; and 1,830 m at 44°–44°60'.

For statistical analyses, we used BMDP (Dixon et al. 1981) and the Statistical Package for the

TABLE 2.—Coefficients of correlation (*r*) between stream habitat or geomorphic variables and trout standing stock in Wyoming streams. Coefficients are shown only if they are significant at $P \leq 0.10$.

Variable	Forest streams (N=65)	Rangeland streams (N=26)
Stream measurements		
Average reach width	-0.42 ^a	-0.52 ^a
Average reach velocity		-0.52 ^a
Width: depth ratio	-0.46 ^a	-0.48
Rubble substrate	-0.24	-0.48
Gravel substrate	0.22	
Silt-sand substrate		0.37
Reach gradient	-0.17	
Geomorphic measurements		
Reach elevation	-0.20 ^a	0.52
Midrange basin elevation	-0.41 ^a	
Stream order		-0.57 ^a
Basin area		-0.28
Basin relief	-0.37 ^a	-0.66 ^a
Relief ratio	-0.35 ^a	-0.40 ^a
Channel slope		-0.37
Drainage density	-0.22	

^a Based on a log₁₀ transformation of the independent variable.

Social Sciences (SPSS; Nie et al. 1975). Correlation analysis was used to determine the correlation (and its significance) between each independent variable and trout standing stock, as well as relations between geomorphic and stream habitat variables.

The relations between those geomorphic and stream habitat variables that were significantly correlated ($P \leq 0.10$) to trout standing stock were investigated further by canonical correlation. If a pair of stream habitat or a pair of geomorphic variables were highly multicollinear ($R \geq 0.75$), one of the two was excluded from analysis to eliminate redundant variables. The remaining variables were used to generate a canonical model for both forest and rangeland streams. Canonical correlation coefficients (R_c) were computed such that the linear combination of stream habitat variables (variate *u*) was maximally correlated to the linear combination of geomorphic variables (variate *v*). Canonical models enable the investigation of more than one relation between the variable sets because they are generated independently (Levine 1977).

Normal probability plots and standardized residual plots were inspected to detect violations of regression assumptions and to determine if logarithmic transformations of certain independent variables were valid (Zar 1974). If logarithmic transformations increased the variance accounted

TABLE 3.—Correlation coefficients between stream habitat and geomorphic variables that were significantly correlated with trout standing stock.

Stream habitat variables	Geomorphic variables									
	Reach elevation	Mid-range basin elevation	Basin relief	Relief ratio	Drainage density	Stream length	Stream order	Basin area	Basin perimeter	Channel slope
Forest streams (critical $r = 0.21$; $N = 65$; $P \leq 0.10$)										
Average reach width		0.36	0.56		-0.28					
Rubble substrate	0.29	0.43	0.28		-0.22					
Gravel substrate	-0.47	-0.45		0.28						
Reach gradient					0.54					
Width: depth ratio	0.41	0.44			0.28					
λ Rangeland streams (critical $r = 0.33$; $N = 26$; $P \leq 0.10$) λ										
Average reach width	-0.67		0.62			0.63	0.62	0.60	0.55	
Average reach velocity	-0.60		0.56			0.52	0.71	0.55	0.46	
Rubble substrate	-0.44		0.53	0.46			0.50			0.58
Silt-sand substrate			0.35			-0.48		-0.44	-0.48	
Width: depth ratio	-0.50						0.33			
	λ		λ			λ	λ	λ	λ	

for by at least 5%, the transformed variable was chosen over the untransformed one for inclusion in multiple-regression analyses.

Variables significantly correlated ($P \leq 0.10$) with trout standing stock were analyzed further for their combined influences on trout standing stock by means of BMDP all-subsets, multiple linear regression (Dixon et al. 1981). This program was used to generate a series of regression models. Each model was then evaluated to determine if the variables included in the model related to trout standing stock in a way that was consistent with results of other studies and current biological thought. After "nonsense" models were excluded, the model with the highest adjusted coefficient of determination, R_a^2 , was chosen (Neter and Wasserman 1974). When the models were tested against an independent data set, prediction error was computed as the difference between predicted and actual standing stock divided by the predicted value and expressed as a percentage.

Results

Data from a total of 91 stream reaches were analyzed, 38 from file information and 53 from our own sampling; 65 were in forests and 26 in rangelands. Many of the variables analyzed in both forest and rangeland streams were significantly ($P \leq 0.10$) correlated with trout standing stock. In forest streams, five of the 10 stream variables and five of the 11 geomorphic variables were significantly correlated with trout standing stock (Ta-

ble 2); in rangeland streams, five of the 10 stream variables and six of the 11 geomorphic variables were significantly correlated with trout standing stock.

Upon inspection of plots of each independent variable and trout standing stock, three independent variables suggested a curvilinear relation with standing stock. These three variables were rated from 1 (low standing-stock range) to 3 (high standing-stock range) to yield a more linear relation between the rated predictor variable and trout standing stock:

reach elevation (RE),

1 = $RE < 2,150$ m,

2 = $RE > 2,355$ m,

3 = $2,150 \leq RE \leq 2,355$ m;

midrange basin elevation (MRE),

1 = $MRE \leq 2,000$ m or

$MRE \geq 2,600$ m,

2 = $2,000 < MRE < 2,325$ m or

$2,475 < MRE < 2,600$ m,

3 = $2,325 \leq MRE \leq 2,475$ m;

width: depth ratio (WD),

1 = $WD \leq 10$ or $WD \geq 33$,

2 = $23 \leq WD \leq 32$,

3 = $11 \leq WD \leq 22$.

Analysis demonstrated a significant relation between trout standing stock and rated reach elevation ($r_a^2 = 0.25$) in forest streams and between midrange basin elevation ($r_a^2 = 0.40$) and width: depth ratio ($r_a^2 = 0.45$) in rangeland streams.

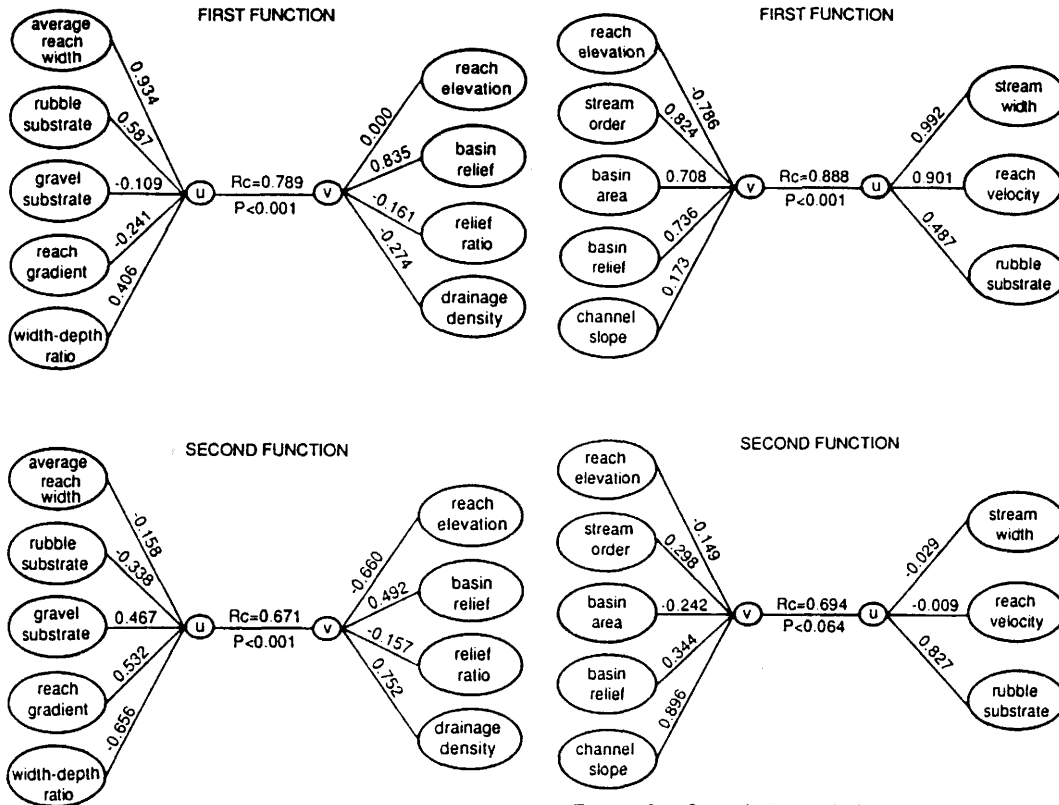


FIGURE 1.—Canonical correlation (R_c) between measures of stream habitat (canonical variate u) and drainage basin geomorphology (canonical variate v) in forest streams.

FIGURE 2.—Canonical correlation (R_c) between measures of stream habitat (canonical variate u) and drainage basin geomorphology (canonical variate v) in rangeland streams.

Relations between Stream Habitat and Geomorphic Variables

Stream habitat and geomorphic variables significantly correlated with trout standing stock were analyzed further to determine their relations with each other. For forest streams, 14 statistically significant correlations existed between the five stream habitat variables and the five geomorphic variables (Table 3). Each of the geomorphic variables correlated significantly with one to four stream habitat variables. In rangeland streams, 23 statistically significant correlations were observed between the five stream habitat variables and the eight geomorphic variables (Table 3).

Canonical analysis indicated significant correlations between the stream habitat variate and geomorphic variate. In forest streams the canonical correlation (Figure 1) between the canonical variates for the first function was highly significant ($R_c = 0.79$; $P \leq 0.001$). Canonical variate v re-

flected drainage basin size whereas canonical variate u reflected stream size. The second function also was highly significant ($R_c = 0.67$; $P \leq 0.001$). Canonical variate v reflected drainage basin features that indicated decreased discharge response time to rainfall events, and canonical variate u reflected stream channel adjustments to this decreased response time.

A significant relation between the canonical variates for the first function also was found in rangeland streams ($R_c = 0.89$; $P \leq 0.001$; Figure 2). Again, canonical variate v reflected drainage basin size and canonical variate u reflected stream size. The second function was also significant ($R_c = 0.69$; $P = 0.064$). Canonical variate v reflected basin gradient, and canonical variate u may have reflected food-producing areas and instream cover.

Regression Models

Three multiple-regression equations describing trout standing stock as kilograms/hectare (Y) were

TABLE 4.—Correlation coefficients and mean prediction errors for tests of models relating trout standing stock in Wyoming streams to stream habitat and geomorphic variables.

Statistic	Com- bined model	Stream model	Geomor- phic model
Forest (N = 11)			
Correlation coefficient (<i>r</i>)	0.80	0.32 ^a	0.75
Prediction error (%)	73	103	101
Rangeland (N = 8)			
Correlation coefficient (<i>r</i>)	0.96	0.90	0.69
Prediction error (%)	18	29	50

^a Not significant ($P > 0.05$). All other *r* values in this table are significant.

developed for both forest ($N = 65$) and rangeland ($N = 26$) streams. In one model (combined model), both geomorphic and stream habitat variables were used (Table 3); the second and third models incorporated only stream habitat or geomorphic variables, respectively. Geomorphic variables dominated the combined models for both forest and rangeland streams: three of the four independent variables were geomorphic measures. Similar relations were observed between stream habitat models and geomorphic models for both forest and rangeland streams, but different independent variables were incorporated.

Forest stream models variously included three stream and four geomorphic variables: average reach width (*ARW*), width:depth ratio (*WD*), and gradient (*G*), and rated reach elevation (*RRE*), relief ratio (*RR*), drainage density (*DD*), and basin relief (*BR*).

Combined variables: *Flow*

$$Y = 447.8 + 67.5RRE - 153.7 \log_{10}(RR + 1) - 35.7DD - 263.1 \log_{10}(ARW + 1);$$

Basin Relief

$$R_a^2 = 0.51; P \leq 0.001. \quad \text{Width}$$

Stream variables:

$$Y = 408.2 - 189.7 \log_{10}(ARW + 1) - 113.9 \log_{10}(WD + 1) - 12.4G;$$

$$R_a^2 = 0.31; P \leq 0.001.$$

Geomorphic variables:

$$Y = 471.5 + 99.4RRE - 138.2 \log_{10}(BR + 1) - 123.6 \log_{10}(RR + 1);$$

Relief

$$R_a^2 = 0.36; P \leq 0.001.$$

Rangeland stream models included two stream and four geomorphic variables: rated width:depth ratio (*RWS*) and average reach velocity (*ARV*)

and rated midrange basin elevation (*RMRE*), basin perimeter (*BP*), channel slope (*CS*), and basin relief (*BR*).

Combined variables: *Flow*, *Perimeter*

$$Y = 200.3 + 36.1RMRE - 0.85BP - 138.7 \log_{10}(CS + 1) + 50.5RWD;$$

$$R_a^2 = 0.64; P \leq 0.001.$$

Stream variables: *W/D*, *Velocity*

$$Y = 39.2 + 71.0RWD - 197.1ARV;$$

$$R_a^2 = 0.57; P \leq 0.001.$$

Geomorphic variables:

$$Y = 487.6 + 53.3RMRE - 160.1 \log_{10}(BR + 1);$$

Basin Relief

$$R_a^2 = 0.52; P \leq 0.001.$$

Following development of the regression equations, an independent data set was obtained from Wyoming Game and Fish Department and Wyoming Water Research Center records to test each model. In all models except the forest stream model, a significant correlation between model predictions and actual standing stock was obtained (Table 4). Rangeland stream tests yielded higher correlations and lower prediction errors than those associated with forest streams when computations followed Binns and Eiserman (1979). The combined model for each stream type gave the best test results.

Discussion

Relations between measures of drainage basin geomorphology, stream habitat quality, and trout standing stock were demonstrated in this study by the numerous univariate correlations between geomorphic and stream habitat variables, the high canonical correlations between geomorphic variates and stream habitat variates, and the extent to which geomorphic variables accounted for variance in the standing stock of trout. Platts (1979) and Parsons et al. (1981) also looked at the relations between drainage basin geomorphology and stream habitat. Platts (1979) found that as stream order increased, stream width, depth, and the percent of rubble substrate also increased, whereas the percent of pool habitats, channel gradient, and the percent of gravel substrate decreased. Parsons et al. (1981) correlated a habitat condition score generated from measured features of stream habitat to four measures of drainage basin geomorphology. All of these relations combine to provide substantial evidence that stream habitat is a func-

tion of geologic processes within the drainage basin.

Geomorphic variables dominated (three of four variables) our multiple-regression models where both variable types were incorporated. In addition, when used separately, trout standing stock was predicted as accurately with geomorphic variables as it was with stream habitat variables. Other studies have successfully used measures of drainage basin geomorphology to predict salmonid standing stock or abundance in streams (Ziemer 1971; Burton and Wesche 1974; Swanston et al. 1977). These observations suggest that geomorphic variables are useful in predicting the potential habitat quality of trout streams.

Our data confirm that small, gently sloping drainage basins produce the best trout habitat. Basin relief, relief ratio, and gradient indicate (by their negative relation to trout standing stock) that a large drop in elevation over the drainage basin leads to reduced trout habitat quality. Branson et al. (1981) stated that high basin relief resulted in greater channel slope and increased drainage density, both of which were negatively related to trout standing stock in our study. The combined effect of watershed features, such as increased basin slope (basin relief and relief ratio), increased channel slope (gradient), and a more dendritic drainage pattern (drainage density), may tend to decrease response time of stream discharge to rainfall events. Drainage basins with these characteristics, when subjected to high-intensity thunderstorms (which are common in Wyoming), generally have greater flow variability, decreased storage of water in depressions and as groundwater, and lower base flows (Viessman et al. 1977). Low base flows and high flow variability result in poor habitat quality for trout (Binns and Eiserman 1979; Wesche et al., in press).

Highest trout biomass was associated with the transition zone between forest and rangeland stream types, which occurred between elevations of 2,100 and 2,355 m in forest streams and 2,100 and 2,224 m in rangeland streams. Platts (1979) found a similar situation in Idaho, and Elser (1968) observed the best habitat quality at the transition between high-gradient, boulder-substrate habitat (characteristic of forest streams) and lower-gradient, gravel-substrate habitat (characteristic of rangeland streams).

Increasing stream size, as reflected by geomorphic variables, resulted in reduced trout density in our study. This relation may be the result of a decrease in relative abundance of riparian cover

or an increase in human impact with increasing stream size. Data presented by Conder (1982) indicated that as stream order increased in the Big-horn Basin of Wyoming human impact on the aquatic and riparian resources increased.

Statistical evidence leads us to the conclusion that the relation between drainage basin geomorphology and trout standing stock is the result of a functional link between measurable features of a drainage basin and stream habitat. This linkage may enable the use of simple measures of drainage basin geomorphology to predict potential habitat quality for trout.

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